

ORIGINAL RESEARCH

Fire history in northern Sierra Nevada mixed conifer forests across a distinct gradient in productivity

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Abstract

Background Understanding the role of fire in forested landscapes is fundamental to fire reintroduction efforts, yet few studies have examined how fre dynamics vary in response to interactions between local conditions, such as soil productivity, and more broadscale changes in climate. In this study, we examined historical fre frequency, seasonality, and spatial patterning in mixed conifer forests across a distinct gradient of soil productivity in the northern Sierra Nevada. We cross-dated 46 diferent wood samples containing 377 fre scars from 6 paired sites, located on and off of ultramafic serpentine soils. Forests on serpentine-derived soils have slower growth rates, lower biomass accumulation, and patchier vegetation than adjacent, non-serpentine sites. Due to these diferences, we hypothesized that historical fre frequency and spatial extent would be reduced in mixed conifer forests growing on serpentine soils.

Results Fire scars revealed a history of frequent fre at all of our sites (median composite interval: 6–22.5 years) despite clear diferences in soil productivity. Fire frequency was slightly shorter in more productive non-serpentine sites, but this diference was not consistently signifcant within our sample pairs. While fres were frequent, both on and off of serpentine, they were also highly asynchronous, and this was largely driven by differing climate–fre relationships. Fires in more productive sites were strongly associated with drought conditions in the year of the fre, while fres in less productive serpentine sites appeared to be more dependent on a cycle of wet and dry conditions in the years preceding the fre. Widespread fres that crossed the boundary between serpentine and nonserpentine were associated with drier than normal years.

Conclusions In our study, fne-scale variation in historical fre regime attributes was linked to both bottom-up and top-down controls. Understanding how these factors interact to create variation in fre frequency, timing, and spatial extent can help managers more efectively defne desired conditions, develop management objectives, and identify management strategies for fre reintroduction and forest restoration projects.

Keywords Dendrochronology, Fire–climate interactions, Fire ecology, Fire return intervals, Fire regime, Fire scars, Soil productivity, Serpentine, Ultramafc

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Resumen

Antecedentes El entender el rol del fuego en paisajes forestales es fundamental en los esfuerzos destinados a su reintroducción; sin embargo, pocos estudios han examinado cómo la dinámica del fuego varia en respuesta a las interacciones entre condiciones locales, como la productividad del suelo, y cambios a gran escala en el clima. En este estudio, examinamos la frecuencia histórica del fuego, su estacionalidad, y patrones espaciales, en bosques mixtos de coníferas a lo largo de un gradiente de distinta productividad del suelo en el norte de la Sierra Nevada. Para ello, cruzamos datos de 46 muestras de leños conteniendo 377 cicatrices de fuego en seis sitios apareados ubicados dentro y fuera de suelos serpentinos ultramórfcos. Los bosques en los suelos derivados de serpentinas tuvieron una tasa de crecimiento menor, menos acumulación de biomasa y vegetación en parches, que los suelos de sitos adyacentes no serpentinos. Debido a esas diferencias, hipotetizamos que la frecuencia histórica y la extensión espacial de los incendios puede ser más reducida en bosques mixtos de coníferas que crecen en suelos serpentinos.

Resultados Las cicatrices revelaron una historia de fuegos frecuentes en todos los sitios (el intervalo de la mediana compuesta fue de entre 6 y 22,5 años), a pesar de las claras diferencias en la productividad de los suelos. La frecuencia de fuegos fue ligeramente más corta en sitios no serpentinos más productivos, aunque esta diferencia no fue consistentemente signifcativa dentro de los pares analizados. Cuando los fuegos fueron frecuentes, tanto en suelos serpentinos como en no serpentinos, se presentaron también como altamente asincrónicos, y esto fue motivado mayoritariamente por diferencias en la relación entre fuego y clima. Los fuegos en los sitios más productivos estuvieron fuertemente asociados con las condiciones de sequía en el año de ocurrencia del fuego, mientras que los fuegos en sitios serpentinos menos productivos parecen ser más dependientes de ciclos de condiciones de sequía y humedad en los años precedentes a los fuegos. Los fuegos que se expandieron y cruzaron los bordes entre suelos serpentinos y no serpentinos se asociaron con períodos más secos que en años normales.

Conclusiones En nuestro estudio, la variación a escala fna en los atributos de los regímenes históricos de fuego, estuvo ligada a las variaciones de control de los tipos *bottom-up*y *top-down*. El entender cómo esos factores interactúan para crear variaciones en la frecuencia del fuego, el tiempo de ocurrencia, y su extensión espacial, puede ayudar a los gestores a defnir de una manera más efectiva las condiciones deseadas, desarrollar objetivos de manejo, e identifcar estrategias para la reintroducción del fuego y en proyectos de restauración.

Introduction

Understanding the factors that infuence and control fre regime attributes is fundamental to fre reintroduction efforts in montane forests in the western USA. Land managers and scientists often use fre regime classifcations as a way to broadly characterize and communicate predominant patterns in fre frequency, seasonality, and extent within a specifc vegetation type or area. Yet, these characteristics can vary widely within individual forest stands and across landscapes, infuenced by local site conditions (e.g., topography, soils, fuel availability, ignitions) and broadscale changes in climate (Beaty and Taylor 2001, Swetnam and Baison [2003\)](#page-19-0).

Bottom-up controls, like topography, vegetation type, and disturbance history exert a strong infuence on fre regime characteristics by impacting the amount, continuity, and availability of fuels (Taylor [2000;](#page-19-1) Heyerdahl et al. [2001](#page-18-0); Stephens [2001](#page-18-1); Taylor and Skinner [2003](#page-19-2); Beaty and Taylor [2008](#page-17-0); Gill and Taylor [2009;](#page-17-1) Ireland et al. [2012](#page-18-2)). For example, variation in slope and aspect can infuence fre patterns by afecting vegetation type and abundance, fuel moisture, and fammability and by creating barriers to fre spread (Agee [1993;](#page-17-2) Taylor and Skinner [2003](#page-19-2); Beaty and Taylor [2008](#page-17-0)). Basic physiological diferences among

diferent compositions of tree species (e.g., needle type and litter production) can infuence fuel fammability and fre spread, resulting in stark diferences in fre frequence and severity among forest types (Taylor [2000](#page-19-1); Gill and Taylor [2009](#page-17-1)). Mosaics of past fres and forest treatments can limit the extent and severity of subsequent fres by reducing the abundance and availability of fuels to burn (Collins et al. [2009](#page-17-3); Ireland et al. [2012](#page-18-2); Stephens et al. [2023\)](#page-18-3). Bottom-up controls can also interact, with one another as well as with broader regional factors (i.e., climate, regional drought), resulting in variation in fre occurrence and spread at both local and landscape scales (Skinner et al. [2009](#page-18-4)).

Site productivity (defned as the capacity to accumulate biomass) plays an important role in determining the amount, continuity, and availability of fuel to burn. However, the infuence of productivity on fre regime attributes is not well defned, particularly in montane forests. Theory predicts that more frequent or intense disturbances may be required in highly productive ecosystems to minimize competition and maximize species diversity, while less frequent or lower intensity disturbance may be sufficient in sites with lower productivity (Huston [1979;](#page-18-5) Kondoh [2001\)](#page-18-6). A positive relationship between

productivity and disturbance has been documented in Mediterranean shrublands, with less productive sites experiencing fewer and less severe fres than more productive sites (Saford and Harrison [2004](#page-18-7); Pausas and Bradstock [2007](#page-18-8)). This link between fire and productivity has been attributed to fuel production and continuity, with faster accumulation of biomass on more productive sites supporting shorter intervals between fres (Saford and Harrison [2004\)](#page-18-7).

Climate is one of the primary top-down drivers of fre occurrence and spatial patterns in western forests (Swetnam and Baisan [2003](#page-19-0); Westerling et al. [2003;](#page-19-3) Collins et al. [2006](#page-17-4)). Climatic variability, which can often be driven by distant teleconnections such as El Nino-Southern Oscillation, infuences regional moisture and temperature patterns and has been linked to widespread fre occurrence (Swetnam and Baisan [2003\)](#page-19-0). Fire and climate reconstructions in dry pine forests of the southwestern USA have documented a canonical fre–climate relationship with wet years linked to subsequent dry years with more fre (Swetnam and Betancourt [1998](#page-19-4); Brown et al. [2008](#page-17-5); Roos et al. [2022](#page-18-9)). In these xeric forest types, antecedent wet conditions were considered necessary to produce suffcient fuels, while dry conditions in the year of the fre were critical for fire ignition and spread. This wet-thendry pattern has also been associated with widespread fre years in the mixed conifer forests of the Sierra Nevada (Stephens and Collins [2004;](#page-18-10) Taylor and Beaty [2005](#page-19-5); Taylor and Scholl [2012](#page-19-6)); however, it does not appear as consistent across individual study areas as it is in the southwestern USA. Several other historical fre reconstructions in the Sierra Nevada have only found strong linkages to dry conditions in the year of the fre (Swetnam and Baisan [2003](#page-19-0); Moody et al. [2006](#page-18-11); Gill and Tay-lor [2009](#page-17-1)). These inconsistent fire–climate relationships could be due to difering understory fuel types (grass/ herbaceous vs. shrub/woody), for which lighter fuel types would be more responsive to the wet then dry cycles (Collins et al. [2006](#page-17-4)). Another factor contributing to these inconsistencies across regions/study areas is overall site productivity. Increased antecedent moisture may not be as important in more productive forest types because there is generally sufficient fuel to support fire spread (Heyerdahl et al. [2001\)](#page-18-0).

Forests growing on distinctly diferent soils that also have old trees and remnant tree material provide a unique opportunity to examine the relative importance of top-down and bottom-up controls on historical fre regimes. In the Sierra Nevada of California, mixed conifer forests occur across a broad gradient of elevation, aspect, and soil types, including less productive serpentine soils. Serpentine soils are notable for their low levels of essential nutrients (calcium, potassium, nitrogen, phosphorous), high concentrations of heavy metals (magnesium and iron), and toxic trace elements (chromium, nickel, cobalt) (Kruckeburg [1984;](#page-18-12) Alexander [2007\)](#page-17-6). Compared to more productive sites, mixed conifer forests growing on serpentine soils are generally characterized by slower growth rates, lower biomass accumulation, and more open and variable vegetation structure (Kruckeburg [1984](#page-18-12); Saford and Harrison [2008](#page-18-13); Saford and Mallek [2010](#page-18-14); DeSiervo et al. [2015\)](#page-17-7). Serpentine sites tend to favor certain tree and understory species over others (Kruckeburg [1984\)](#page-18-12), which vary in litter compaction and drying rates (Banwell et al. [2013\)](#page-17-8) as well as fammability (Fonda et al. [1998](#page-17-9); Stevens et al. [2020](#page-19-7)). Nutrient-limited serpentine soils are also less likely to experience rapid shifts in plant composition in response to variation in climate (Briles et al. [2011](#page-17-10)). In California, serpentine outcrops range in size from a few square meters to hundreds of square kilometers (Grace et al. [2007](#page-17-11)), resulting in numerous places where serpentine soils contact more productive non-serpentine soils. Most of the research in these transitional zones has focused on quantifying changes in abiotic or biotic factors, while the infuence that these abrupt changes in soil have on fuels and disturbance processes like fre has received very little scientifc attention (Boyd et al. [2009\)](#page-17-12).

In this study, we examine fre regime attributes in mixed conifer forests growing across a gradient of soil productivity. Our study area encompasses the contact zone between serpentine and non-serpentine soils, which can have considerable variation in soil chemistry and plant productivity (Boyd et al. [2009](#page-17-12)). To identify diferences in soil productivity within our study area, we frst measured soil nutrient composition within our sampling sites. We hypothesized that serpentine soils would have substantially lower concentrations of essential soil nutrients (i.e., calcium, potassium, nitrogen, phosphorous) and higher concentrations of heavy metals (magnesium) than adjacent non-serpentine sites, even when sites were situated in close proximity. We then used fre scar analysis to evaluate diferences in fre regime attributes (frequency, seasonality, and spatial patterns) between sites with difering levels of soil fertility. We hypothesized that areas with lower soil fertility, which would historically have been characterized by lower biomass, slower forest growth rates, and more patchy vegetation (Kruckeburg [1984](#page-18-12); Saford and Harrison [2008](#page-18-13); Saford and Mallek [2010](#page-18-14)), would have experienced less frequent and more variable fres.

Methods

Study area

Our study sites are located near the community of Meadow Valley on the Plumas National Forest in the

northern Sierra Nevada of California (Fig. [1](#page-3-0)). The climate is characterized as Mediterranean, with a warm and dry summer period that extends into the fall. Most of the precipitation in the study area falls during the winter months, with mean annual precipitation totaling approximately 1060 mm (Abatzoglou [2013](#page-17-13)). Mean monthly temperatures range from a minimum of 2.2 °C in December to a maximum of 20.3 °C in July (Abatzoglou et al. [2009\)](#page-17-14).

The study sites are situated on the western edge of some of the most extensive ultramafc terrain in the northern Sierra Nevada (Fig. [1](#page-3-0)). Soils derived from

ultramafc rocks, broadly referred to as serpentine soils, are characterized by their unique chemistry—specifcally high levels of iron and magnesium and relatively low levels of calcium (Alexander [2007\)](#page-17-6). Within our study area, soils range from well-drained, productive Alfsols (suborder: Xeralfs) to less productive, strongly leached Ultisols (suborder: Xerults) (USDA Natural Resources Conservation Service [2023](#page-19-8)).

The vegetation in our study sites is predominantly conifer forest with varying mixtures of tree species. Stands growing on non-serpentine soils have a wider diversity

Fig. 1 Fire scar sample sites within three paired sites on the Plumas National Forest in northern California

of tree species, including a mixture of ponderosa pine (*Pinus ponderosa* Dougl. ex Laws.), sugar pine (*Pinus lambertiana* Dougl.), white fr (*Abies concolor* [Gordon & Glend.] Lindl. ex Hildebr.), incense cedar (*Calocedrus decurrens* [Torr.] Florin), Douglas fr (*Pseudotsuga menziesii* [Mirb.] Franco), and California black oak (*Quercus kelloggii* Newb.). Many of these species are absent in adjacent forest stands growing on serpentine soils, which are primarily composed of Jefrey pine (*Pinus jefreyi* Grev. & Balf.), incense cedar, and Douglas fr. Although forest structure has changed dramatically over the past century due to a combination of fre exclusion and past harvest, compositional diferences between forests on and off serpentine remain apparent (Fig. [2,](#page-4-0) Briles et al. [2011](#page-17-10)). Structural diferences are also evident with mixed conifer forests on serpentine soils generally having fewer trees, lower biomass, and lower surface fuels than adjacent non-serpentine forests (DeSiervo et al. [2015\)](#page-17-7).

The landscape within our study area has been managed in both the distant and recent past. The Mountain Maidu have lived in this area for thousands of years and used fre to enhance young shoot growth for basket weaving, to clear hunting grounds, enhance California black oak stands, and reduce fre risk (Dixon [1905](#page-17-15); Young [2003](#page-19-9); Stephens et al. [2023](#page-18-3)). More recently, forest thinning occurred in the area between 2003 and 2006, and small non-commercial trees and associated fuels were piled and burned (USDA Forest Service [2020](#page-19-10)). No wildfres have been recorded within the boundary of our study since 1900, when written records began.

Our study included three pairs of sites, each containing one sampling site established on serpentine soils and one adjacent sampling site on non-serpentine soils (Fig. [1](#page-3-0)). Sites were all located within an area of approximately 2 km², ranging in elevation from 1090 to 1220 m. The distance between sampling sites within a pair ranged from 22 m to over 800 m and the area sampled within each site ranged from 0.6 to 7.8 ha (Table [1](#page-4-1)). Paired sites were selected to represent similar aspects, topographic positions, and elevations.

Field sampling

In June 2020, we surveyed each of the six sites for frescarred trees, stumps, and logs. Samples were collected opportunistically to maximize the completeness of fre dates and were based on the presence of multiple wellpreserved fre scars (Van Horne and Fulé [2006\)](#page-19-11). During sample collection, Global Positioning System (GPS) coordinates were recorded for each sample and subsequently used to defne the boundary of each sampling site.

We used a chainsaw to cut wedges from 57 fre-scarred trees, stumps, and logs. We focused our sampling on

Fig. 2 Example of contemporary forest conditions in (**a**) serpentine and (**b**) non-serpentine sites within a pair

ponderosa pine (37% of samples), Jeffery pine (22%), and incense cedar (41%). Samples were sanded and polished to a high sheen and cross-dated using standard dendrochronological techniques (Stokes and Smiley [1977\)](#page-19-12).

To evaluate diferences in productivity between study sites, we collected soil samples from fve plots (0.04 ha) that were randomly placed within each of our six sampling sites (30 plots total). At each plot, we collected approximately 150 g of soil at 5–15 cm depth, from four locations 12 m from plot center and combined them to create one sample per plot. Soil samples were air dried, sieved, and analyzed for organic matter (OM), cation exchange capacity (CEC), pH, nitrate (NO3-N), phosphorous (Olsen-P), potassium (K), calcium (Ca), magnesium (Mg), and sodium (Na) by the UC Davis Analytical Laboratory. At the same four sampling locations, the depth of the litter and duf was measured and combined to obtain an average for each plot.

Analysis

Soil productivity

To examine variation in soil nutrients across our study area we used non-metric multidimensional scaling (NMDS) ordination utilizing a Bray–Curtis dissimilarity index with the metaMDS function in the vegan package in R (Oksanen et al. 2022). Correlation coefficients of soil nutrients with both axis 1 and axis 2 scores were used to identify which nutrients were most responsible for the variation described in the ordination. We used a multivariate analysis of variance (ANOVA) procedure for distance matrices (ADONIS) based on 9999 restricted permutations of the data to compare soil characteristics among serpentine and non-serpentine sites. We also used ANOVA to test for diferences in soil nutrients and litter and duf depth between serpentine and non-serpentine sites. Two variables, soil Ca and the ratio of Ca to Mg, were log transformed to meet the assumptions of the model. All statistical analyses were conducted in R (R Core Team [2019\)](#page-18-16).

Fire seasonality

The intra-ring position of each fire scar was used to assign a calendar year to each tree ring and fre event in our sample. Each scar was assigned to the early-earlywood, middle-earlywood, late-earlywood, latewood, or ring boundary (dormant season) position (Caprio and Swetnam [1995](#page-17-16)). In this region, earlywood scars are generally considered to be fres that burned in spring and early summer, while latewood scars are classifed as midto late-summer fres, and dormant (ring-boundary) scars are characterized as fres occurring in late summer or fall, when tree growth has ceased for the season (Stephens et al. [2018](#page-18-17)). Scars that could not have their season identifed were classifed as unknown.

Fire frequency

After the initial processing of fre scar data using the FXH2 software package (Grissino-Mayer [2001\)](#page-18-18), we used the R package *burnr* (Malevich et al. [2018](#page-18-19)) to read and process fre scar samples across all sites and to develop fre chronologies for each site. To account for variability in fre frequency at diferent scales, and to facilitate comparisons with other fre history studies, we estimated fre return interval (FRI) using three diferent methods.

We frst calculated the mean number of years between successive fres for individual trees, not including the origin-to-frst scar interval. We then averaged the individual tree FRIs, which were weighted by their corresponding number of intervals, to obtain a point estimate of FRI for each site (Baker and Ehle [2001](#page-17-17)). Tree-based point estimates often yield longer FRIs due to the fact that individual fres may be missed in the fre-scar record of an individual tree. While trees are often charred by lowintensity surface fres, they are not always scarred due to a wide range of factors, including low fuel loads at the base of the tree, wind shifts, high fuel moisture, as well as variations in species-specifc vulnerability to scarring and the lack of previous scars or wounds (Baker and Ehle [2001](#page-17-17); Stephens et al. [2010](#page-18-20)). Given that the absence of a scar on an individual tree does not defnitively indicate the absence of fre in that location for a given year, this metric can be seen as a more conservative estimate of fre extent (Swetnam et al. [2011\)](#page-19-13).

To obtain a more comprehensive record of past fres, we also calculated composite fre histories for each site using (1) all fire-scarred trees $(C00)$ and (2) a filter $(C25)$ that required a minimum of (a) 25% recording trees scarred or (b) at least two fre-scarred trees for each fre year (whichever was smaller) (Dieterich [1980](#page-17-18)). The use of these flters reduces the likelihood of registering a tree scarred by means other than fre but may cause small fres with scars registered by a single tree to be missed. With each fltering method, we calculated the mean, median, minimum, and maximum fre return interval for each sampled site. We also calculated the Weibull median probability interval (WMPI), which represents the median value of a Weibull distribution. This flexible distribution is frequently used to represent the skewed distribution of fre return interval data (Grissino-Mayer [1999](#page-17-19)). We performed one-sample Kolmogorov–Smirnov (K-S) tests to evaluate goodness-of-ft between fre return interval data and the Weibull distribution at each site.

Our fre interval data were not normally distributed. To account for this, we used the non-parametric Kruskal– Wallis test to determine if FRI was diferent between the

paired serpentine and non-serpentine types and among the six diferent sites in our study area. When signifcant diferences were found, we used a pairwise Wilcoxon rank sum test to identify which group or groups varied from the others. We used a two-sample K–S test to compare the distribution of composite FRIs between serpentine and non-serpentine types.

Synchrony and patchiness

To evaluate the relative spatial extent of past fres, we compared the synchrony of fres between sites. To test this, we calculated Jaccard similarity (JS) as

$$
JS = \frac{F_{11}}{F_{11} + F_{10} + F_{01}}
$$

where F_{11} is the number of years in which two sites recorded a fire and F_{10} and F_{01} is the number of years when only one of the two sites recorded a fire. We used linear regression to evaluate the relationship between Jaccard similarity in the fre year and the geographic distance between each of the sites. Analysis of variance was used to investigate the amount of similarity in fre years between sites that occurred within and outside of study pairs. Jaccard similarity values were arcsine transformed to meet model assumptions. We limited individual analyses to the time period when sites had at least one recording tree and a minimum of one fre-scarred tree.

We estimated the degree of potential fre patchiness within a site by calculating the proportion of trees scarred (out of the total number of trees recording) in each fre year, recognizing that trees can also experience fire without scar formation. This is therefore a better indicator of relative fre patchiness, rather than actual fre patchiness, between sites. We tested for diferences in percent scarred between non-serpentine and serpentine types, and among sites and pairs, using the non-parametric Kruskal–Wallis test.

Fire–climate relationships

In order to examine the relationship between fre occurrence and historical climate patterns, we applied a superposed epoch analysis (SEA) of the composite fre history dataset (Swetnam and Baisan [1996](#page-19-14); Grissino-Mayer and Swetnam [2000](#page-18-21)) using the *burnr* package in R (Malevich et al. [2018\)](#page-18-19). Climate data consisted of a reconstructed Palmer Drought Severity Index (PDSI) for our sampling area from 1632 to 1885 (the time period where trees were recording in at least two of the three pairs) and was extracted using the North American Drought Atlas web application (Cook et al. [2010](#page-17-20)). PDSI is a widely used index that integrates temperature, precipitation, and soil moisture to estimate drought severity (Palmer [1965](#page-18-22)). SEA isolates reconstructed PDSI values for each fre event

year (year=0), as well as the years prior to (lag years−1 to−6) and following (lags 1 to 4) the fre event. It calculates mean PDSI values for each time step and then assesses their statistical signifcance by comparing them to bootstrapped values from a random set of fre years (Grissino-Mayer and Swetnam [2000](#page-18-21)). To investigate the infuence of climate on fre extent, we applied the SEA using fres that were limited to a single site and fres that burned more than one site. We also conducted separate SEA for serpentine and non-serpentine sites.

Results

Soil productivity

The NMDS ordination of soil nutrients resulted in a two-dimensional solution with a fnal stress of 0.017 and revealed a distinct clustering of serpentine and nonserpentine plots, driven almost entirely by the frst axis (Fig. [3\)](#page-7-0). Multivariate ANOVA on the underlying distance matrix indicated that this clustering was signifcant (pseudo- $F_{1,28}$ =97.2, $p=0.001$). The variables with the strongest correlation with NMDS axis 1 were concentrations of K (R^2 =0.95), Ca (R^2 =0.80) and the ratio of Ca:Mg $(R^2=0.81)$.

Serpentine sites had signifcantly higher concentrations of $NO₃$, Mg, and pH, and lower concentrations of K and Ca than non-serpentine sites in all pairs ($p \le 0.001$). The ratio of Ca:Mg, a strong proxy for productivity (DeSiervo et al. [2015](#page-17-7)), was eight times greater on non-serpentine (5.1 ± 1.9) than serpentine sites (0.6 ± 0.6) . Non-serpentine sites also had deeper litter and duf than serpentine sites $(F_{1,24} = 27.5, p < 0.001)$.

Fire seasonality

We were able to estimate fre seasonality on 229 (61%) of the fre scars in our samples. Excluding unknown fre seasons, the highest proportion of scars were found in the latewood (37%) and dormant (39%) positions, indicating that most fires occurred after trees stopped growing. The remaining fre scars were found in the early-earlywood (2%), middle-earlywood (9%), and late-earlywood (13%) positions. The seasonality of fires in non-serpentine samples was very similar to serpentine samples, with almost equivalent proportions of earlywood fres (25% vs. 24%) and only slight diferences in the proportion of fres documented in the latewood and dormant positions (Fig. [4](#page-7-1)).

Fire frequency

We cross-dated samples from a total of 46 trees from old stumps, logs, and standing snags within our six sampling sites. A total of 377 fre scars were identifed, registering 94 diferent fre years in which at least one tree was scarred (Fig. [5](#page-8-0)). Fires spanned a period of over 500 years, with the earliest recorded fre occurring in 1402 and the

Fig. 3 Nonmetric multidimensional scaling (NMDS) ordination of soil nutrients from 30 plots on serpentine and non-serpentine sites. Arrows represent soil variables that were signifcantly correlated with NMDS axis 1 or 2. The length and direction of the arrows correspond to the magnitude and direction of correlation with the NMDS axes. Bold values had the highest correlation with NMDS axis 1

Fig. 4 Proportion of scars in the earlywood, latewood, and dormant (ring-boundary) position in non-serpentine (*n*=130 scars) and serpentine (*n*=99 scars) sites

latest in 1918. Most samples had a limited number of scars prior to the 1700s, and only one fre was recorded by one sample after 1893. The mean minimum age of sampled trees was estimated to be 258 years (range: 80–641; SD: 111 years); however, actual tree ages were underestimated because many samples did not reach pith due to wood decay.

We defned our period of analysis as 1632 to 1885, based on the criteria of having at least one tree recording at each site in two of our three pairs. These dates were selected to allow for statistical comparisons of fre intervals between paired serpentine and non-serpentine sites as well as among pairs. Within this time

Fig. 5 Fire chronology for the three paired sites in our study area. Horizonal lines represent individual trees in serpentine (gray) and adjacent non-serpentine (black) sites. Solid lines indicate the period of time when a tree was recording (i.e., after it had been scarred at least once). Each vertical tick is a dated fre scar

period, we documented 88 diferent fre years in which at least one tree was scarred.

Fire was a frequent occurrence at all of our sites prior to the 1900s. Within our analysis period (1632– 1885), estimates of mean and median FRI varied slightly depending on the composite criteria used. When all fire-scarred trees were used (i.e., no filter— C00), mean FRI ranged from 4.8 to 12 years, median FRI ranged from 4.5 to 12 years, and the Weibull median probability interval (WMPI) ranged from 4.4

to 11 years (Table [2](#page-8-1)). Composite FRIs lengthened when a 25% filter (C25) was applied, which is to be expected with the exclusion of isolated fire events that only scar a single tree. Mean FRI ranged from 7.1 to 25.1 years, median FRI ranged from 6 to 22.5 years, and WMPI ranged from 6.8 to 23.7 years (Table [2\)](#page-8-1). Approximately 63% $(n=46)$ of intervals were 10 years or less and 25% ($n=18$) were 5 years or less. Approximately 86% of the intervals ($n=63$) in our composite fire history were 20 years or less. The intervals had a similar

Table 2 Composite fire return interval (FRI) estimates from paired serpentine and non-serpentine sites. The analysis period for both estimate methods was 1632–1885. Superscript letters indicate signifcant diferences between sites (i.e., a is signifcantly diferent than b, c). *WMPI* Weibull median probability interval

Pair	Site	Composite FRI (no filter)				Composite FRI (25% filter) ^a			
		Mean	Median	WMPI	Range	Mean	Median	WMPI	Range
	Non-serpentine 1	4.8 a	4.5	4.4	$1 - 15$	7.1 ^a	6	6.8	$4 - 17$
	Serpentine 1	7.5 ^{ab}	6	6.1	$1 - 23$	7.4^{ab}	8		$1 - 14$
2	Non-serpentine 2	6.4^{a}	5	4.9	-37	7.8 ^{ab}	6	6.9	$2 - 22$
	Serpentine 2	9.6 ^{bc}	8	9	$4 - 26$	11.8 ^{bd}	9	11.2	$6 - 26$
3	Non-serpentine 3	11^{bc}	9	10	$2 - 25$	17.9 ^{cd}	16	16.1	$3 - 47$
	Serpentine 3	12 ^c	\mathcal{P}		$2 - 29$	25.1°	22.5	23.7	$7 - 47$

^a Composite criteria: fires burned two or more trees or > 25% of recording trees

distribution on serpentine and non-serpentine sites (Kolmogorov–Smirnov, $D=0.28$, $p=0.07$), with more than half of the intervals recorded as 10 years or less $(Fig. 6)$ $(Fig. 6)$ $(Fig. 6)$.

As expected, tree-based point FRIs were slightly longer than our composite FRIs. Mean FRI averaged 9.3 to 26.1 years across sites and median FRI ranged from 9.5 to 24.6 years (Table 3). We documented an average of eight scars per tree (range: 3–19). Approximately 41% $(n=134)$ of the intervals recorded on individual trees were 10 years or less and 10% (*n*=33) were 5 years or less. Four of the samples recorded intervals of 2 years (2 intervals) and 3 years (4 intervals); all of these trees were on non-serpentine soils and three of the trees occurred in pair 2. An estimated 73% of the fire return intervals (*n*=238) recorded on individual trees were 20 years or less.

Mean FRI trended slightly longer in serpentine sites compared to non-serpentine sites, but this difference was not consistently significant within our paired sites (Table [2](#page-8-1), Fig. [7](#page-10-0)). When all fire-scarred trees were used to calculate the composite (C00), mean FRI in serpentine was significantly longer than non-serpentine, but only for the sites in pair 2 (Table [2](#page-8-1)). We found no significant difference in mean FRI between any of the paired serpentine and non-serpentine sites using the other methods of calculation (PFRI or C25). We did find differences among the three paired sites. Point and composite mean FRI (C00 and C25) was significantly longer in pair 3 compared to the other two pairs (Kruskal–Wallis, C00: *χ*²=26.0, *P* < 0.001; C25: χ^2 = 22.6, *P* < 0.001).

Table 3 Individual tree fre return interval (Point FRI) estimates from paired serpentine and non-serpentine sites. The analysis period was 1632–1885. Superscript letters indicate signifcant diferences between sites (i.e., a is signifcantly diferent than b, c)

^a Average of individual tree Point FRIs, weighted by their corresponding number of intervals

Synchrony and patchiness

Of the 88 discrete fre years recorded by fre scars within our six sites between 1632 and 1885, 53 (60%) were recorded on only a single site, and 29 (33%) were recorded by only a single tree at a single site (Fig. [8.](#page-11-0)). Approximately two-thirds of these (35 fres) were recorded in a single non-serpentine site, while the remaining one-third (18 fres) were recorded in a single serpentine site. Our chronology identifed 35 years that fre was recorded in more than one site. There were 4 years (1729, 1782, 1829, 1843) when fres were recorded in four of the six sites, and a single year (1776) when five of the six sites recorded a fire.

Similarity in fre years was signifcantly and negatively correlated with the distance between sites $(R^2=0.52, p=0.002)$, with sites in close proximity sharing the highest percentage of fre years (Fig. [9](#page-11-1)). We

Fig. 6 The distribution of composite fire return intervals (C25) in serpentine and non-serpentine sites

Fig. 7 Violin plots showing (**a**) point and (**b**) composite (25% flter) estimates of fre return interval in non-serpentine and serpentine sites within pairs. Points represent median values

found some evidence of synchrony in paired serpentine and non-serpentine sites, but this was likely due to the close proximity of these sites within study pairs (Fig. [9\)](#page-11-1). The highest percentage of boundary-spanning fires (35%) was documented in pair 1, where sites were situated less than 25 m apart. The other two pairs shared 17% and 19% of fres, respectively. Fires crossed the boundary between serpentine and non-serpentine sites in 27 of the fre years; however, 10 of these boundaryspanning fres were part of larger fre events that were recorded in three or more sites. The proportion of trees scarred (out of total number recording) in each fre year was highly variable and did not difer signifcantly among sites, pairs, or between serpentine and non-serpentine forest types (Fig. [10\)](#page-12-0).

Fire–climate relationships

Climatic conditions in our study area ranged from severe drought (PDSI:−5.8) to very wet (PDSI: 4.1) over the analysis period (1632–1885). In non-serpentine sites, fires were associated with drought conditions (negative PDSI) in the year of the fire (Fig. [11](#page-13-0), $p < 0.01$). Fires in serpentine sites had a significant association with conditions prior to the fire but showed no relationship with PDSI in the year of the fire. Serpentine sites showed a significant relationship

Fig. 8 The number of sites burned in each fre event between 1632– 1885

with wet conditions (positive PDSI) five years prior to fire $(p < 0.05)$, followed by strong drought conditions four years prior to the fire event $(p < 0.01$, Fig. [11](#page-13-0).).

In years when fires were limited to a single site, we found no significant relationship between fire occurrence and PDSI. In contrast, more extensive fires (i.e., those recorded at more than one site), occurred in years that were drier than expected (Fig. [12](#page-14-0)).

Discussion

Despite clear diferences in soil productivity, we did not detect consistent diferences in mean FRI between serpentine and non-serpentine conifer forests. This similarity in fre frequency suggests that fre spread potential was similar in serpentine and non-serpentine sites, despite distinct diferences in productivity, biomass accumulation, and fuel continuity. However, it is also important to recognize that these fndings could, at least in part, be a function of sampling close to the edge of the serpentine soil contact zone, where fre spread from adjacent more productive forests might be more likely than the interior of large serpentine barrens, which generally have lower tree densities, more exposed rock, and discontinuous cover of grasses and shrubs.

The range of FRIs across our study area is comparable to those reported by other studies conducted in mixed conifer forests of the Sierra Nevada (Table [4](#page-15-0)). For example, Moody et al. [\(2006](#page-18-11)) used similar methods to derive median composite FRIs for non-serpentine mixed conifer forests on the Plumas National Forest that ranged between 5–15 years (using all scars) and 6–16.5 years (25% filter). The composite FRIs for serpentine sites in our study (median: 8–22.5 years) are also similar to those reported by Taylor and Skinner [\(2003\)](#page-19-2) where documented median composite FRIs ranged from 8 to 15 years for mid-elevation Jefrey pine sites growing on serpentine soils in the Klamath Mountains (Skinner et al. [2018](#page-18-23)). It is worth noting that fre-scar-based reconstructions

Fig. 9 Relationship between Jaccard similarity in shared fre events and the geographic distance between sites. Points represent pairwise comparisons between non-serpentine sites (NonSerp), serpentine sites (Serp), and serpentine and non-serpentine sites within a study pair and outside of a study pair

Fig. 10 The number of fres by the percentage of recording trees scarred within each of the six sampling sites. Values were weighted by the number of trees recording at the time of the fre

likely underestimate fre frequency due to a wide range of factors that infuence tree scarring, such as fuel loading around the base of the tree, conditions at the time of the burn (e.g., wind, fuel moisture), topographic position of the tree, and the presence of previous scars or wounds (Baker and Ehle [2001](#page-17-17); Collins and Stephens [2007;](#page-17-21) Farris et al. [2010](#page-17-22); Stephens et al. [2010](#page-18-20)). Sampling intensity and sampling area can also infuence FRI estimates, with more extensive sampling resulting in shorter composite FRIs as more fre events are captured (Stephens and Collins [2004](#page-18-10); Baker and Ehle [2001\)](#page-17-17). Taking these factors into consideration, historical fre frequency at our sites was likely greater than what we report here (see Table [4\)](#page-15-0).

The highest proportion of scars in our study area were formed by fres that occurred between late-summer and early fall (Stephens et al. [2018](#page-18-17)), when tree growth had ceased for the season, similar to the seasonal timing reported for other studies in northern California (Taylor and Skinner [2003;](#page-19-2) Taylor and Beaty [2005;](#page-19-5) Fry and Stephens [2006](#page-17-23); Moody et al. [2006;](#page-18-11) Gill and Taylor [2009;](#page-17-1) Vaillant and Stephens [2009](#page-19-15); Stephens et al. [2023\)](#page-18-3). The predominance of fires recorded in the latewood and dormant period was consistent among sites and between serpentine and non-serpentine soils (Fig. [4\)](#page-7-1). This timing suggests that many fires in our study area likely occurred in the later part of the dry season and before the beginning of the wet season, as a result of ignitions from either lightning or cultural burning practices, at times when surface fuels were dry enough to promote fre spread (Moody et al. [2006;](#page-18-11) Stephens et al. [2023\)](#page-18-3). The relatively high proportion of fires in

the dormant season also includes the possibility that some Indigenous ignited fres could have occurred in the cooler months, when sunny conditions could have dried out surface fuels (Stephens et al. [2023\)](#page-18-3).

More than half of the fres in our study area were recorded in only a single site, and a third were recorded by only a single tree within a site. Fires that burned beyond the boundary of a single site were most often recorded by sites in close proximity. Considering the relatively small size of our study area (approximately 2 km^2) and the close proximity of our paired sites, these fndings suggest that, while frequent, historical fres may also have been relatively limited in extent. The finding that historical fires were unlikely to burn large portions of the study area in any given year may be due to the abrupt change in parent material and soil type, which have been shown to inhibit fre spread under normal climate conditions (Taylor and Skinner [2003](#page-19-2)). Scattered rock outcrops, perennial streams, and meadows, all of which occur within our study area, can also afect fre spread by breaking up fuel continuity. The frequency of locally limited fires, in combination with variation in topography and abrupt changes in soil type, would have historically resulted in a spatially diverse mosaic of vegetation types, sizes, and age classes (Hessburg et al. [2019\)](#page-18-24). This pattern would have been self-reinforcing, as the patchwork of burns limited the extent of subsequent fres by infuencing the abundance and availability of burnable fuel (Collins et al. [2009;](#page-17-3) Scholl and Taylor [2010;](#page-18-25) Ireland et al. [2012;](#page-18-2) Stephens et al. [2023](#page-18-3)).

Fig. 11 Superposed epoch analysis (SEA) plots for Palmer Drought Severity Index (PDSI) in non-serpentine (top) and serpentine (bottom) sites. Light gray bars indicate wet/cool years and dark gray bars indicate dry/warm years. Dashed line indicates signifcance at the *p*<0.05 level, solid line at the *p*<0.01 level. Asterisk indicates signifcant (*p*<0.05) departure from mean conditions

While it is difficult to distinguish between Indigenous and lightning-ignited fres using fre scars alone (Anderson [2005](#page-17-24); Van Wagtendonk and Cayan [2008](#page-19-16); Roos et al. [2019](#page-18-26)), a few lines of evidence suggest that the frequency of fres in our study area was substantively augmented by Indigenous ignitions. The constant promotion of fire, whether ignited by lightning or humans, has long been a major component of Mountain Maidu land stewardship (Dixon [1905;](#page-17-15) Kroeber [1925](#page-18-27), Stephens et el. [2023](#page-18-3)). Frequent burning, often in relatively small patches, was used to reduce fuels, clear hunting grounds, and promote the growth of plants for food, medicine, and basketry (Dixon [1905](#page-17-15); Kimmerer and Lake [2001](#page-18-28); Stephens et al. [2023](#page-18-3)). In our study, the inclusion of all fre-scarred trees (C00) resulted in the shortest fre return intervals (median FRI: 4.5 to 12 years). Compared to other fltering methods, this method has been identifed as being most efective at capturing smaller, more localized fres that may have been intentionally lit or tended (Roos et al. [2019](#page-18-26)). Most of the fres recorded in our study area also occurred at short intervals of less than 10 years, and many were less than 5 years (Fig. 6 .). This fire interval is difficult to attribute

Fig. 12 Superposed epoch analysis (SEA) plots for Palmer Drought Severity Index (PDSI) for fires that burned a single site (top) and fires that burned more than one site (bottom). Light gray bars indicate wet/cool years and dark gray bars indicate dry/warm years. Dashed line indicates signifcance at the *p*<0.05 level, solid line at the *p*<0.01 level. Asterisk indicates signifcant (*p*<0.05) departure from mean conditions

to lightning fres alone, which would require both frequent strikes and an available fuel bed, the latter of which can be a limiting factor in frequently burned areas or in low-productivity sites where fuel accumulation is slower (Van Wagtendonk and Cayan [2008](#page-19-16); Stephens et al. [2023](#page-18-3)). Unfortunately, we did not have the sampling depth to investigate how land use changes (e.g., the decline in Native American populations after 1850) may have afected fre frequency in our study area. However, several fre history studies have documented a reduction in fre frequency and increase in fre extent following the decline in Native American populations and disruption of cultural burning (Fry and Stephens [2006;](#page-17-23) Gill and Taylor [2009](#page-17-1); Skinner et al. [2009](#page-18-4); Taylor et al. [2016](#page-19-17)).

Fire return intervals in the serpentine and non-serpentine sites in our southern-most study pair were signifcantly longer (median composite FRI: 16–22.5 years) than the FRIs documented in the other two pairs of sites (median composite FRI: 6–9 years; Table [2](#page-8-1)). Although we were not able to explicitly test this, anecdotally, the six **Table 4** Fire return interval estimates from comparable fre history studies in mixed conifer forest in the northern and central Sierra Nevada. *NSN* northern Sierra Nevada, *CSN* central Sierra Nevada, *C00* composite using all fre scars, *C25* composite of fres that scarred 25% of sample, *PFRI* point (individual tree) estimate

^a Values represent the grand mean when FRI estimates were presented for multiple sites

^b FRI calculated for entire study area

sites in our study area occur at relatively similar elevations, slopes, and aspects (see Table [1\)](#page-4-1) and the three pairs were selected to reduce variation in these attributes. One potential explanation for the shorter FRIs in pairs 1 and 2 may be their close proximity to large meadow complexes and creeks. The four trees with extremely short intervals (intervals of 2 and 3 years) were situated less than 200 m from the edge of large meadows. During the drier parts of the year, the continuity of fammable fuels in these habitats could have provided connectivity with other parts of the landscape, collecting fre from adjacent areas and facilitating fire spread to surrounding uplands. The shorter FRIs near the edges of meadows may also have been an indicator of Indigenous burning (Anderson and Carpenter [1991;](#page-17-25) Anderson and Moratto [1996](#page-17-26); Kimmerer and Lake [2001](#page-18-28); Anderson [2005](#page-17-24)). Native Americans in this region used periodic burning in and around meadows to discourage the encroachment of conifers, maintain and promote edible plants, and improve forage for wildlife (Kimmerer and Lake [2001](#page-18-28); Anderson [2005](#page-17-24)).

Climate exerted a strong infuence on fre occurrence and extent in our study area, but the specifc relationship between climatic conditions and fre activity differed between serpentine and non-serpentine sites. Fires in more productive non-serpentine sites were strongly associated with drought conditions in the year of the fre (Fig. [11\)](#page-13-0). This coincides with a number of other studies in northern California mixed conifer forests that have also found that historical fres were correlated with signifcantly drier conditions (Moody et al. [2006;](#page-18-11) Taylor et al. [2008](#page-19-18); Skinner et al. [2009\)](#page-18-4). Drought conditions in the year of the fre would have created highly favorable conditions for fre spread, such as low fuel moisture, high air temperatures, and low relative humidity.

Fire occurrence in our serpentine sites was limited more by preceding climatic conditions than moisture in the year of the fre. In serpentine sites, fres were preceded by a signifcantly wet year, followed by a substantially dry year, 5 and 4 years prior to the fire (Fig. 11). These antecedent wet conditions would have produced more abundant and continuous herbaceous fuels, while a subsequent dry year would have allowed those fuels to cure, increasing the potential for fre spread in future years (Margolis et al. [2022](#page-18-29); Roos et al. [2022](#page-18-9)). While the mechanism behind this wet-then-dry connection to fre occurrence is fairly well established, the longer preceding lag (4–5 years) and lack of a strong fre year signal for our serpentine sites is somewhat puzzling. Skinner et al. ([2008\)](#page-18-30) found a similar 5-year lag with wet conditions leading to more fre occurrence in the Jefrey pine-mixed conifer forests of the northern Baja, Mexico, but this was also connected to a strong drought signal for fre years. Skinner et al. ([2008](#page-18-30)) posited that the longer lag with wet years may have been connected to needle cast, for which the wetter conditions allowed for greater needle production, that was then cast 4–6 years later, resulting in greater fuel continuity. A similar mechanism may be responsible for fre occurrence on our serpentine sites, which are also dominated by Jefrey pine. Alternatively, this longer preceding lag could be connected with understory plant growth, perhaps perennial forbs and shrubs if the understory was

similar to what exists now, which are slower to respond than grasses. Regardless of the specifc mechanism for increased fuel production and continuity, the lack of a drought signal for fre years suggests that these serpentine sites may be somewhat unique in that fammability is not a major limiting factor.

The difference in fire–climate patterns on and off serpentine soils suggests that fre regime attributes in our study area varied in response to interactions between top-down climatic controls and bottom-up site conditions. Soil productivity indirectly afects fre cycles by directly infuencing the amount of biomass that is available to burn (Pausas and Bradstock [2007](#page-18-8)). Compared to more productive sites, forests growing on serpentine soils have substantially lower biomass, slower growth rates, and a more open and heterogenous structure (Kruckeburg [1984](#page-18-12); Saford and Harrison [2008](#page-18-13); Saford and Mallek [2010\)](#page-18-14). However, widely spaced trees in serpentine stands can also increase both light and wind near the soil surface, resulting in faster drying of surface fuels and higher potential for wind-driven fires (Safford and Mallek [2010](#page-18-14); Bigelow and North [2012\)](#page-17-27). These open forests were historically dominated by Jeffery pine, which produces a well-aerated, fammable, long-needle litter layer that is also conducive to fre spread (Fonda et al. [1998\)](#page-17-9). Taken together, these factors suggest that under normal climate conditions, serpentine sites had suitable environmental conditions for fre spread (i.e., higher surface temperatures and wind speeds) but were dependent on canonical patterns of wet and dry years to produce sufficient understory fuels for fres to more easily spread.

We found little overlap in the timing of fres between serpentine and non-serpentine sites, suggesting that fres were often asynchronous. When fres did cross the boundary between serpentine and non-serpentine sites, they were often associated with more extensive fre years (i.e., fres recorded in two or more sites) that occurred during drier than normal conditions. There were 5 years (1729, 1776, 1782, 1829, 1843) when fres were recorded in four or more sites, and all occurred in years with lower PDSI. Several of these years were also identifed as regional fre years by studies synthesizing multiple individual fre history reconstructions (Swetnam and Baisan [2003](#page-19-0); Trouet et al. [2010;](#page-19-19) Taylor et al. [2016](#page-19-17)). According to Taylor et al. [\(2016\)](#page-19-17), 1829 is the most highly recoded fre year in the entire Sierra Nevada, with evidence of burning as far north as southwestern Oregon (Metlen et al. [2018](#page-18-31)). These findings suggest that the abrupt boundary between soil types in our study area may have acted like a flter to fre spread, containing fres under normal climate conditions, while allowing fres to spread across the gradient of soil productivity in drier years (Taylor and Skinner [2003](#page-19-2)).

Conclusion

As the push to increase the pace and scale of fre reintroduction gains momentum, it is important for managers to have reliable information about how fre regime attributes vary in response to local site conditions, as well as more broadscale changes in climate. In our study, fnescale variation in historical fre frequency and timing was linked to both bottom-up and top-down controls, which interacted to infuence the quantity, continuity, and availability of fuels. Fires in our more productive sites were strongly associated with drought conditions in the year of the fre, while fres in our less productive serpentine sites were limited more by fuel amount and availability. In these less productive systems, the capacity for fre spread appeared to be strongly dependent on a cycle of wet and dry conditions in the years preceding the fre to produce sufficient fuels for fires to more readily spread.

The fire scars in our study area recorded a history of frequent, low severity fre at all of our sites and a distinct lack of fre after the beginning of the twentieth century. This dramatic decline in fire frequency has been documented in mixed conifer forests throughout California (Safford and Stevens [2017;](#page-18-32) Bohlman et al. [2021](#page-17-28)) and is likely due to a variety of contributing factors including the removal of Mountain Maidu populations and disruption of cultural practices (Stephens et al. [2023](#page-18-3)), as well as the onset of organized fre suppression on federal lands in the early 1900s (Stephens and Ruth [2005](#page-18-33)). This marked decrease in fre frequency also follows a period of intensive livestock grazing that peaked in the late-1800s that could have broken up the continuity of grass and herbaceous fuels across the landscape and reduced the potential for fre spread (Swetnam and Baisan [2003](#page-19-0); Norman and Taylor [2005](#page-18-34); Taylor et al. [2016](#page-19-17)).

Forest changes in the absence of fre have been well documented for more productive mixed conifer forests in the Sierra Nevada (Parsons and Debenedetti [1979](#page-18-35); Collins et al. [2011](#page-17-29); Knapp et al. [2013](#page-18-36); Stephens et al. [2015](#page-18-37)). Signifcant increases in tree density, shifts in species composition, and reduced structural diversity have occurred at both the stand and landscape scale (Saford and Stevens [2017\)](#page-18-32). Similar changes in forest structure have been documented in fre-excluded forests growing on serpentine soils, but these changes often occur over much longer time periods due to slower growth rates (Sahara et al. 2015). These changes, in combination with the fndings from this study, suggest that frequent, low severity fre was a fundamental component of mixed conifer forests growing on and off of serpentine soils. Understanding how historical fre regime attributes varied in response to both local and broadscale conditions can help managers more efectively defne desired conditions, develop management objectives, and identify

management strategies for fre reintroduction and forest restoration projects in the Sierra Nevada and elsewhere.

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Authors' contributions

MC, SS, BC, and EK designed the study and collected feld samples. CA prepared and dated fre scar samples. MC and HF completed the data analysis. MC wrote the manuscript, with contribution from all authors.

Availability of data and materials

The datasets used during the current study are available from the corresponding author on reasonable request.

Declarations

Competing interests

The authors declare that they have no competing interests.

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